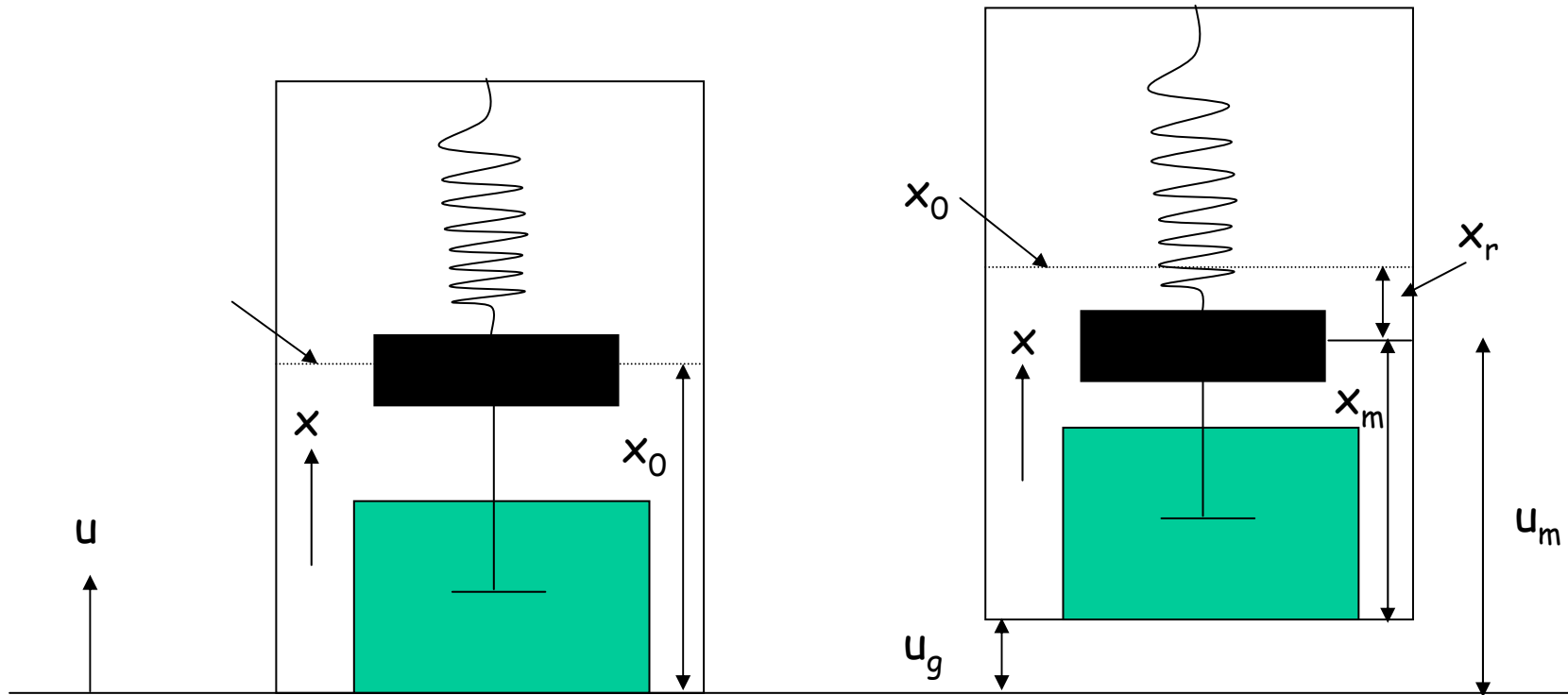
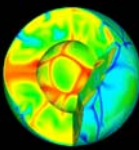




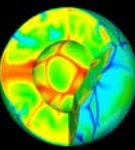
# Seismometer - The basic Principles



$u$  ground displacement  
 $x_r$  displacement of seismometer mass  
 $x_0$  mass equilibrium position



# Seismometer - The basic Principles



The motion of the seismometer mass as a function of the ground displacement is given through a differential equation resulting from the equilibrium of forces (in rest):

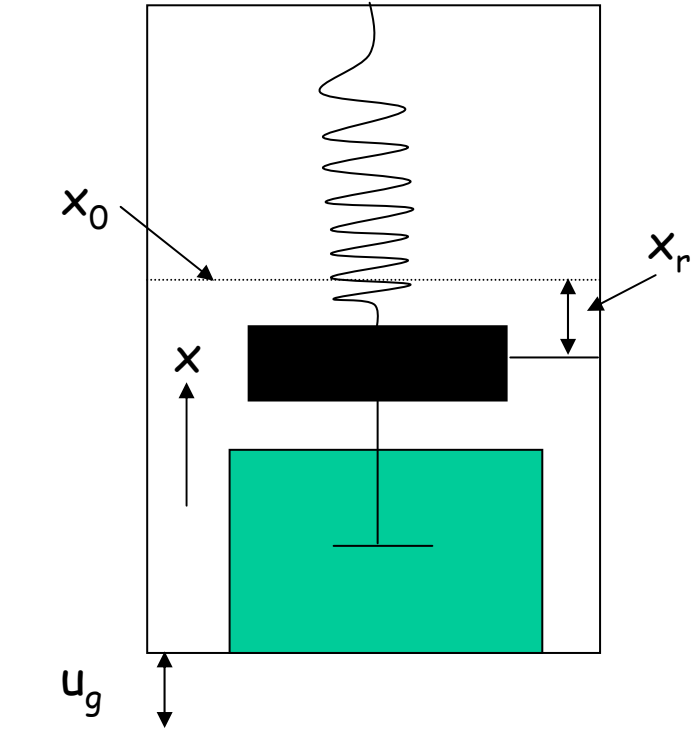
$$F_{\text{spring}} + F_{\text{friction}} + F_{\text{gravity}} = 0$$

for example

$$F_{\text{spring}} = -k x, \text{ k spring constant}$$

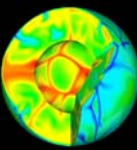
$$F_{\text{friction}} = -D \dot{x}, \text{ D friction coefficient}$$

$$F_{\text{gravity}} = -m\ddot{u}, \text{ m seismometer mass}$$





# Seismometer - The basic Principles



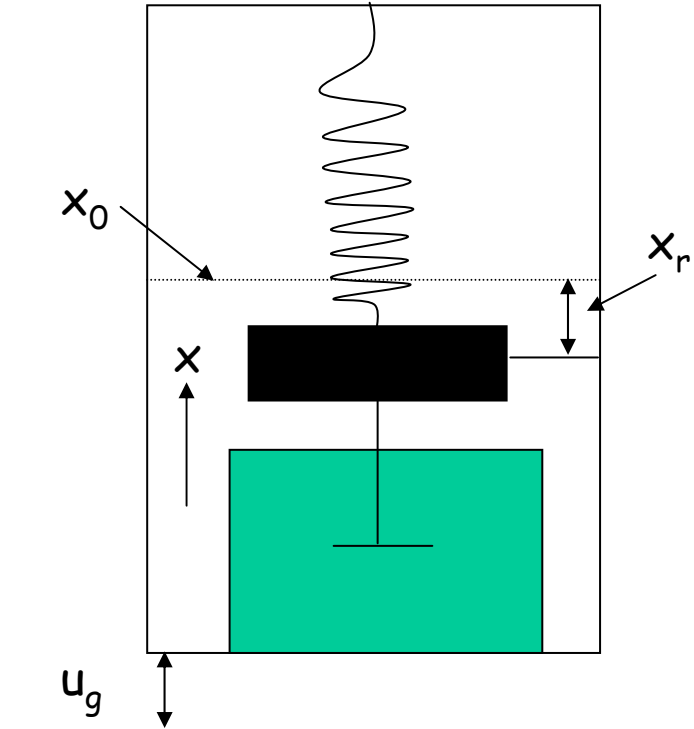
using the notation introduced the equation of motion for the mass is

$$\ddot{x}_r(t) + 2\varepsilon\dot{x}_r(t) + \omega_0^2 x_r(t) = -\ddot{u}_g(t)$$

$$\varepsilon = \frac{D}{2m} = h\omega_0, \quad \omega_0^2 = \frac{k}{m}$$

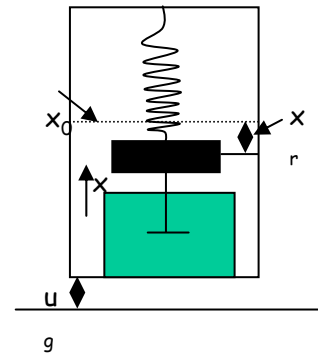
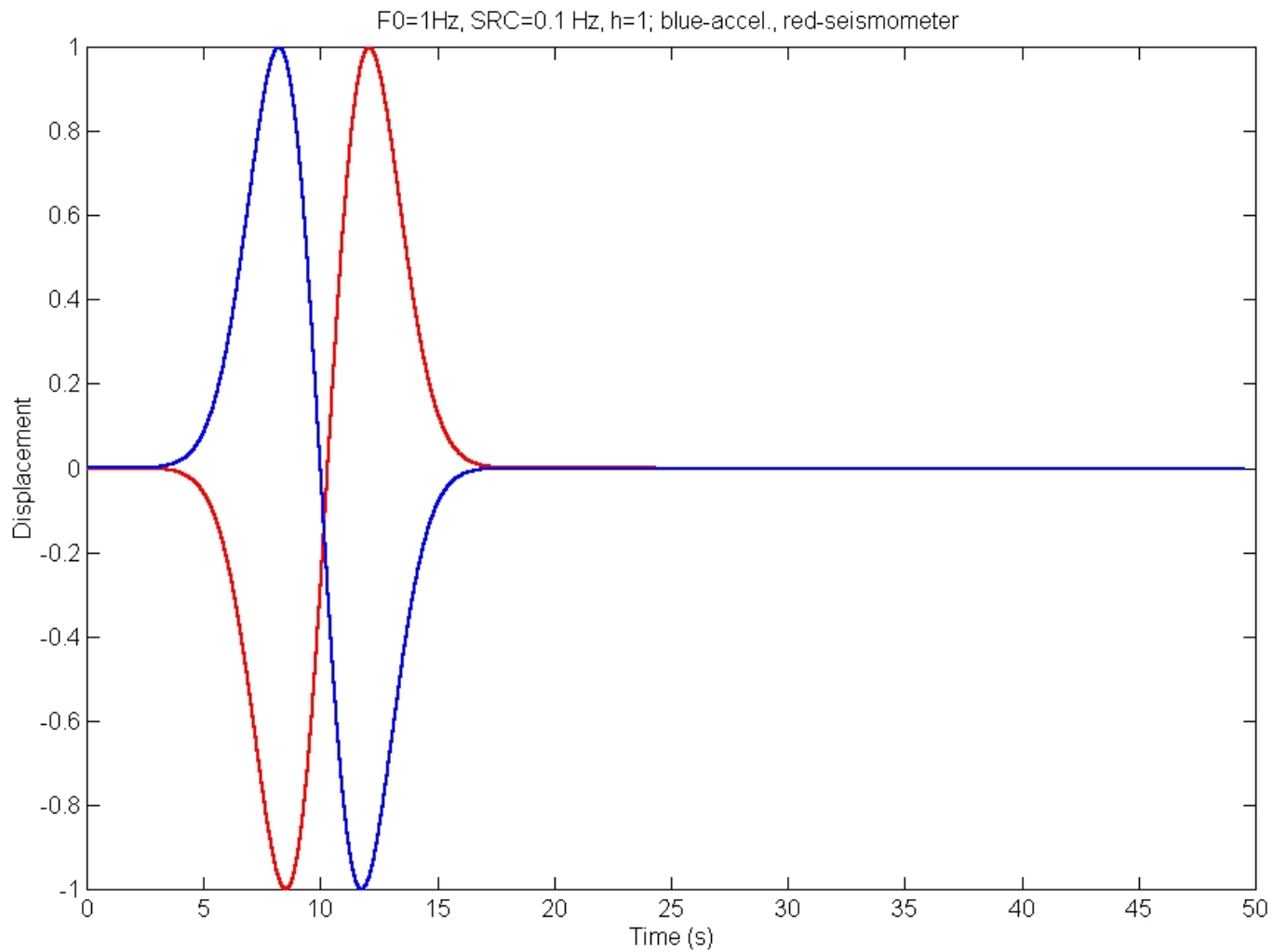
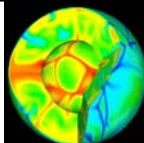
From this we learn that:

- for slow movements the acceleration and velocity becomes negligible, the seismometer records ground acceleration
- for fast movements the acceleration of the mass dominates and the seismometer records ground displacement



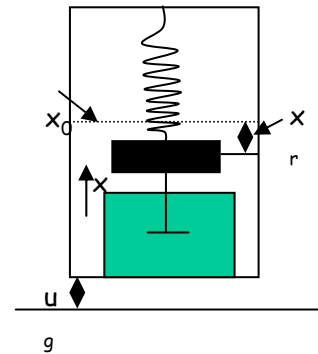
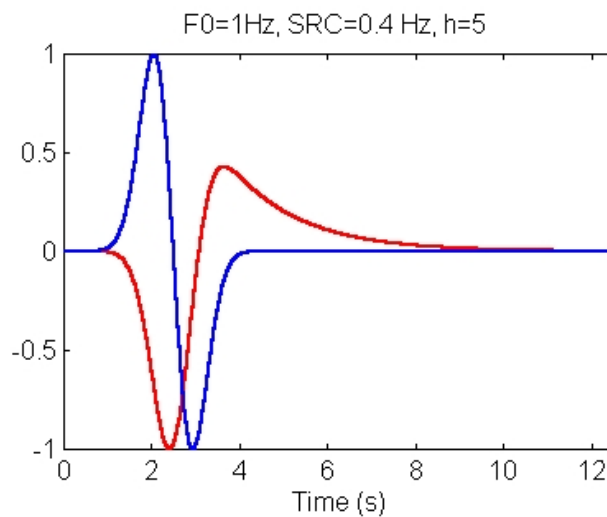
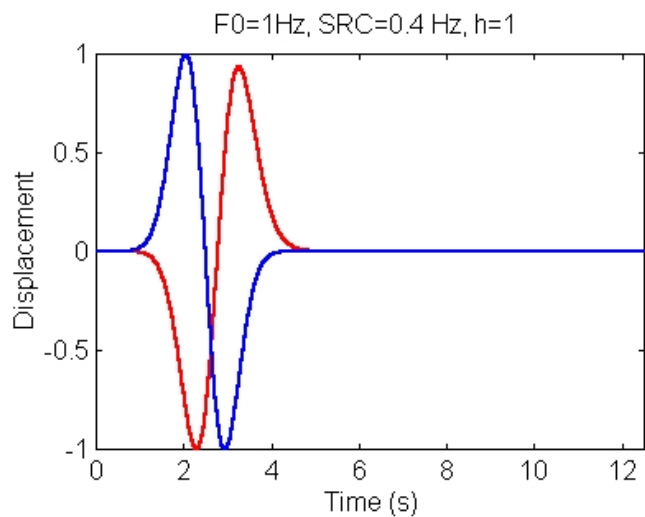
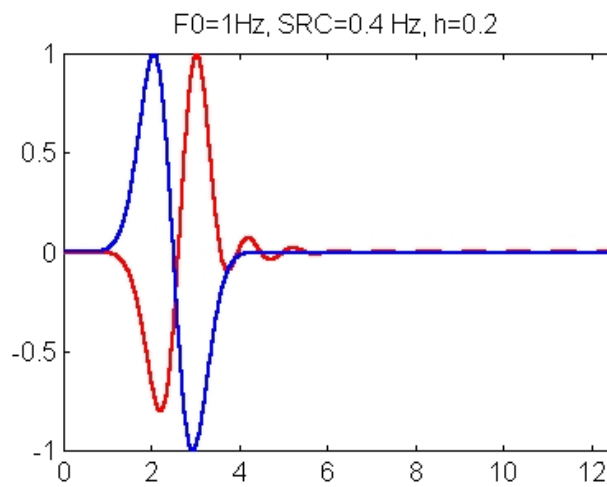
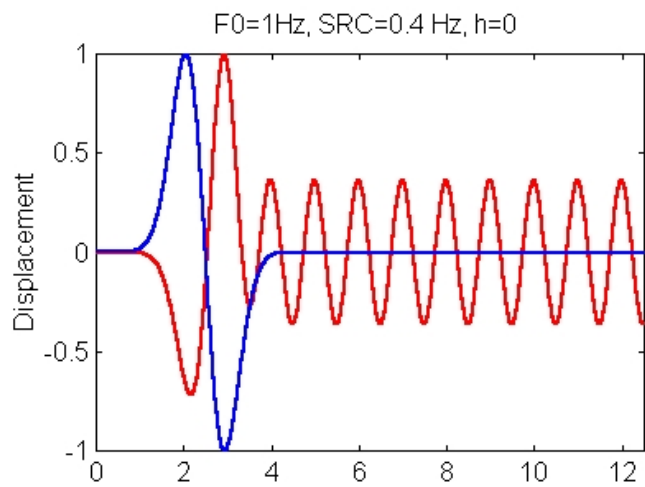
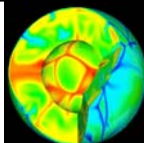


# Seismometer - examples



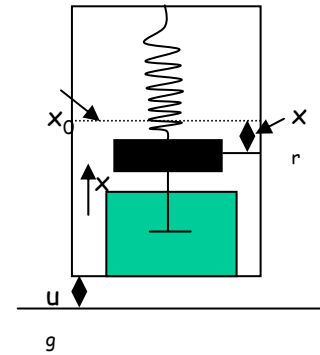
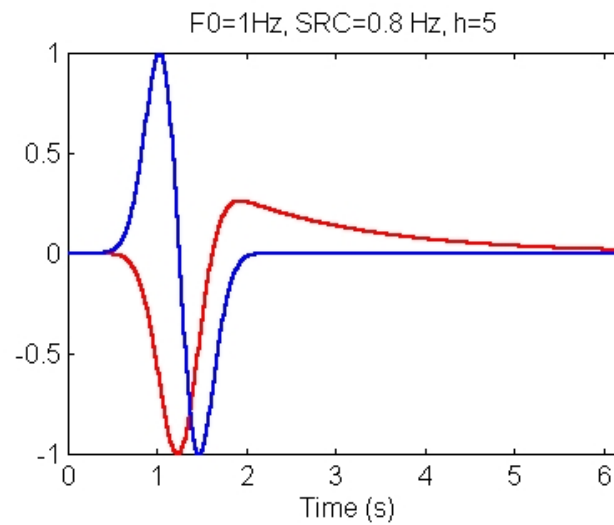
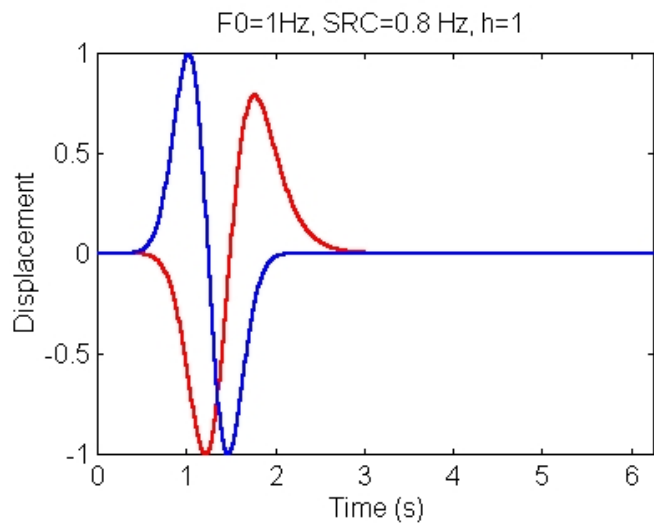
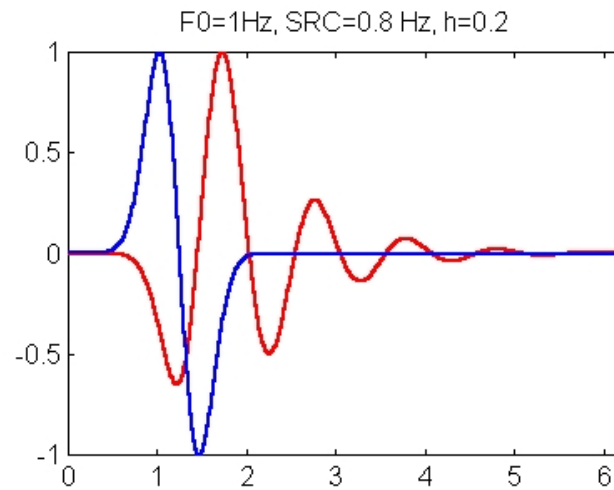
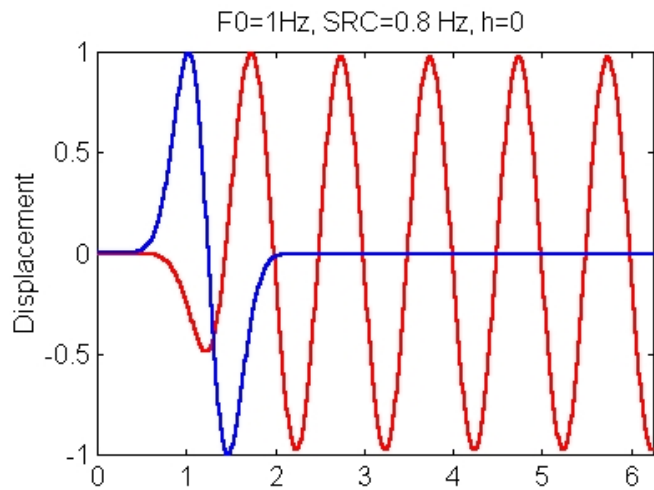
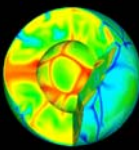


# Seismometer - examples



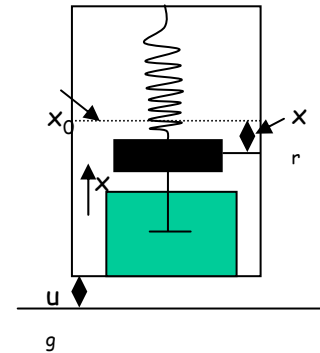
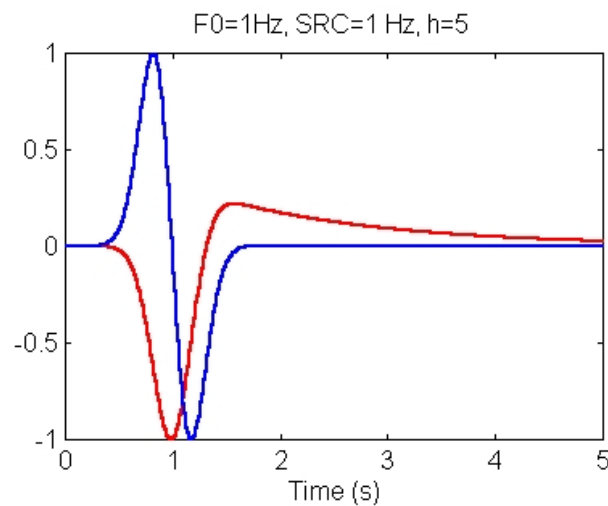
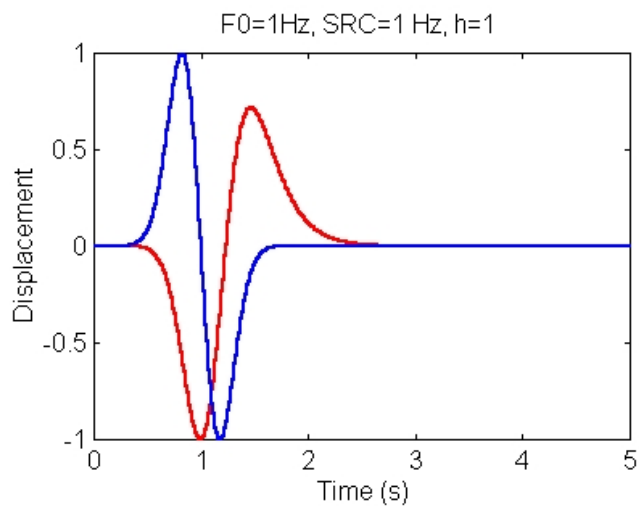
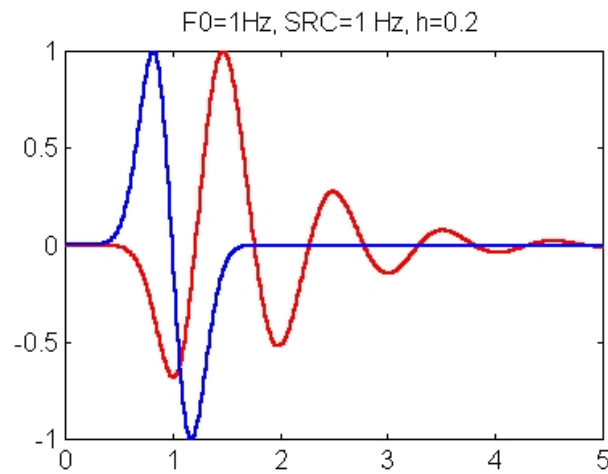
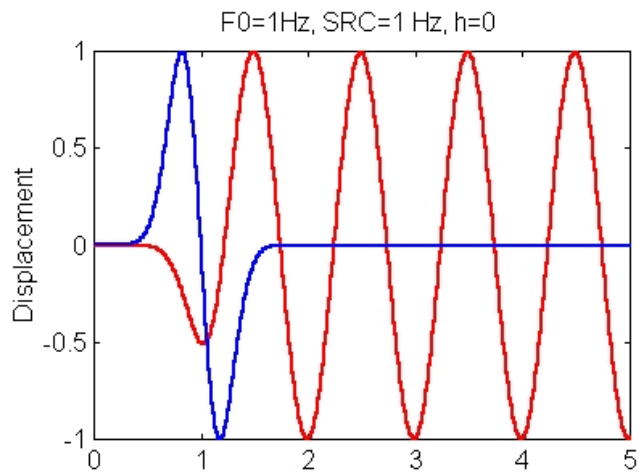
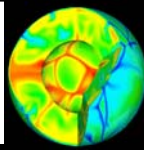


# Seismometer - examples



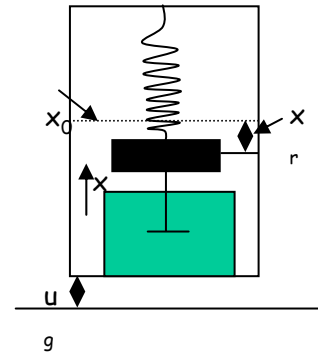
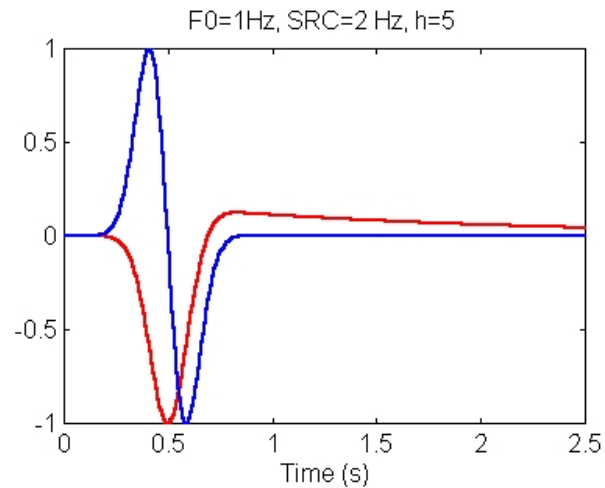
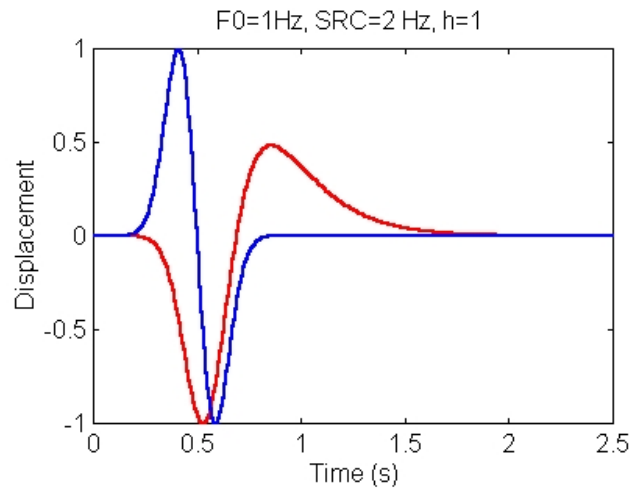
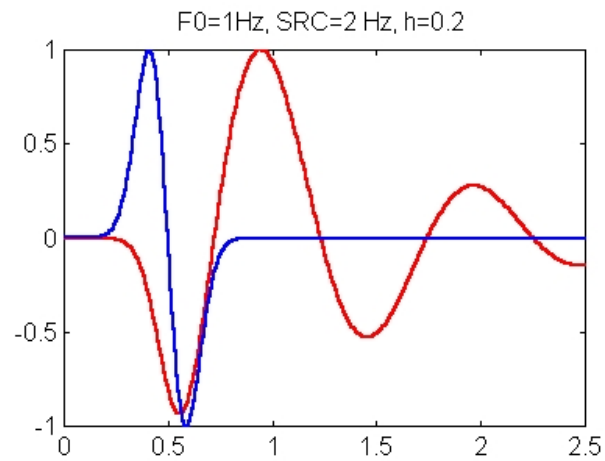
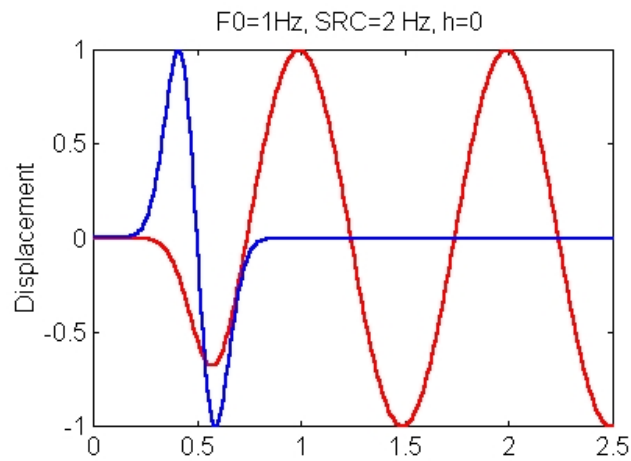
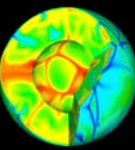


# Seismometer - examples





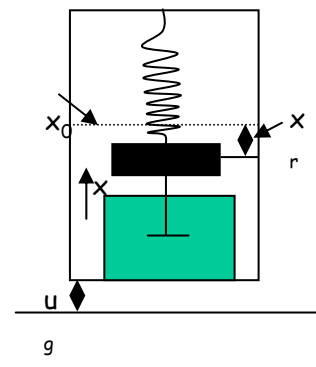
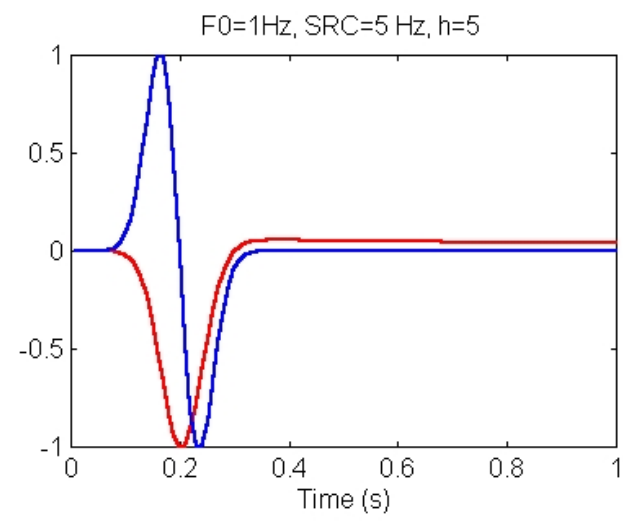
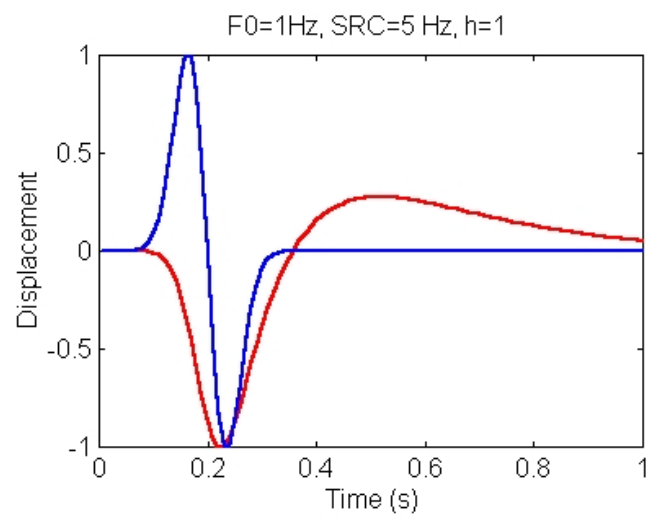
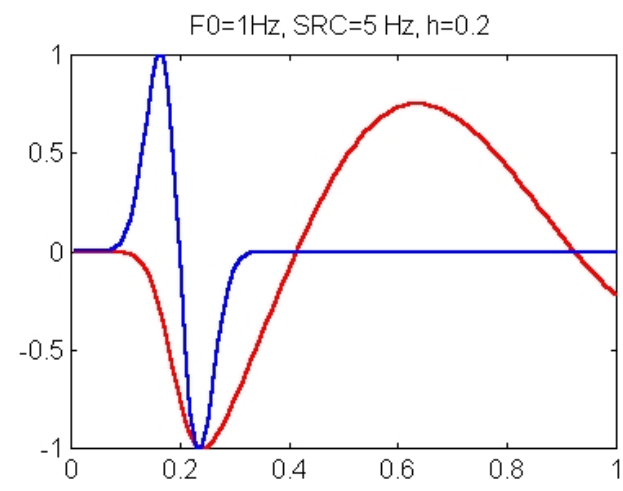
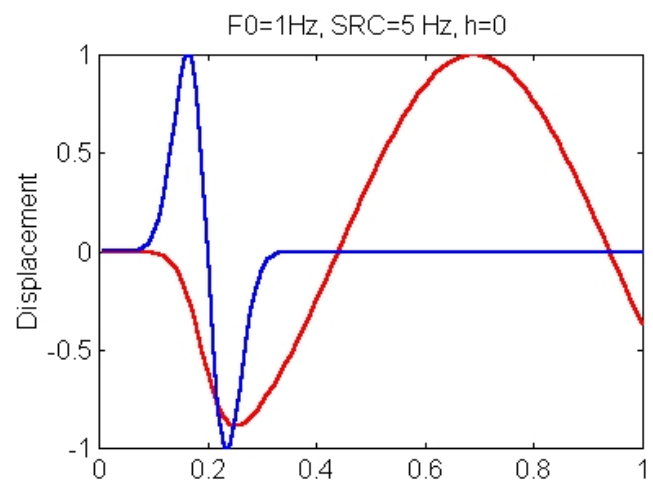
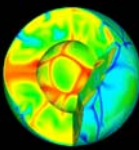
# Seismometer - examples





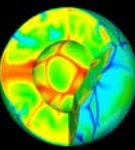


# Seismometer - examples





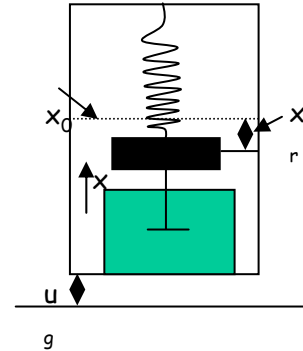
# Seismometer - Questions



1. How can we determine the damping properties from the observed behavior of the seismometer?

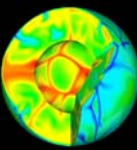
2. How does the seismometer amplify the ground motion? Is this amplification frequency dependent?

We need to answer these question in order to determine what we really want to know:  
The ground motion.





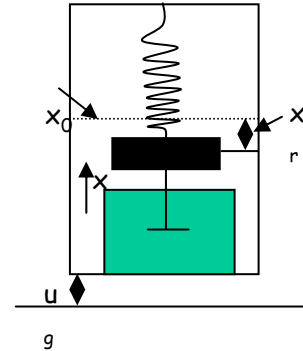
# Seismometer - Release Test



1. How can we determine the damping properties from the observed behavior of the seismometer?

$$\ddot{x}_r(t) + h\omega_0\dot{x}_r(t) + \omega_0^2x_r(t) = 0$$
$$x_r(0) = x_0, \quad \dot{x}_r(0) = 0$$

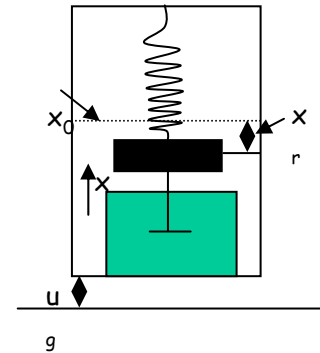
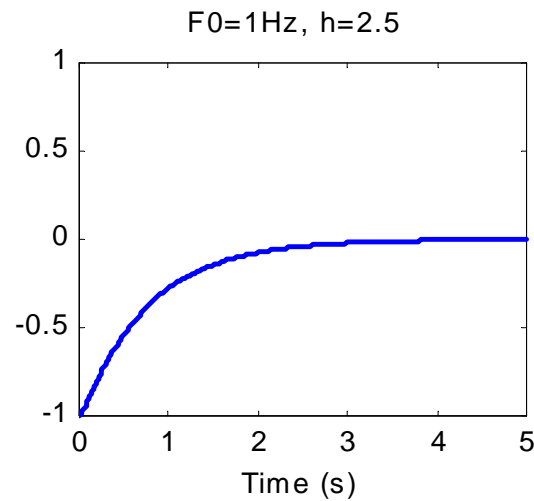
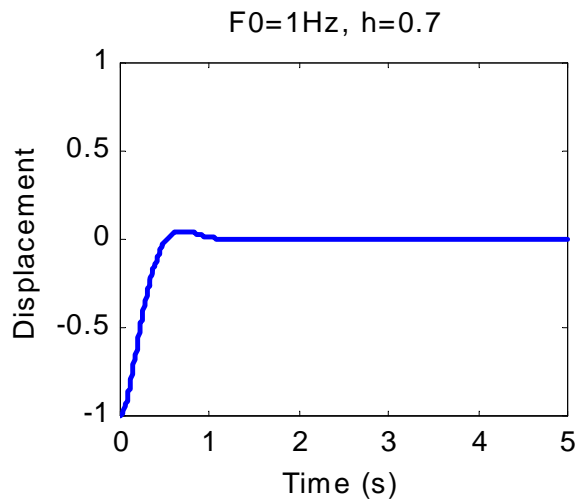
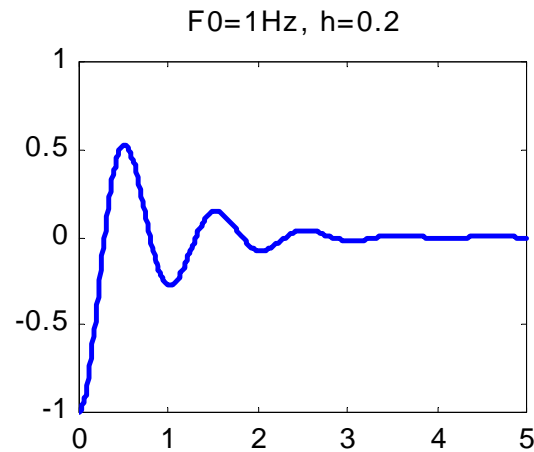
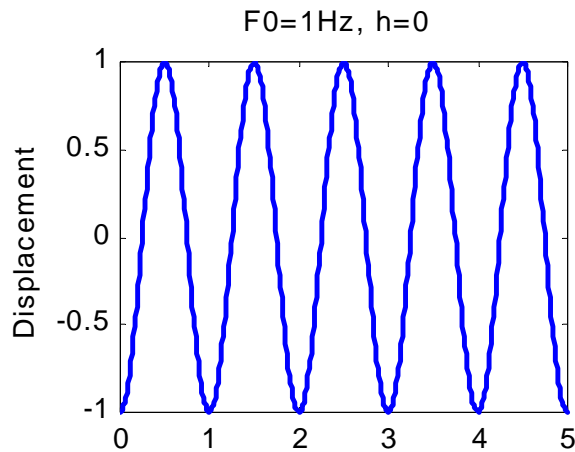
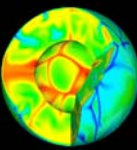
we release the seismometer mass from a given initial position and let it swing. The behavior depends on the relation between the frequency of the spring and the damping parameter. **If the seismometer oscillates, we can determine the damping coefficient  $h$ .**



9

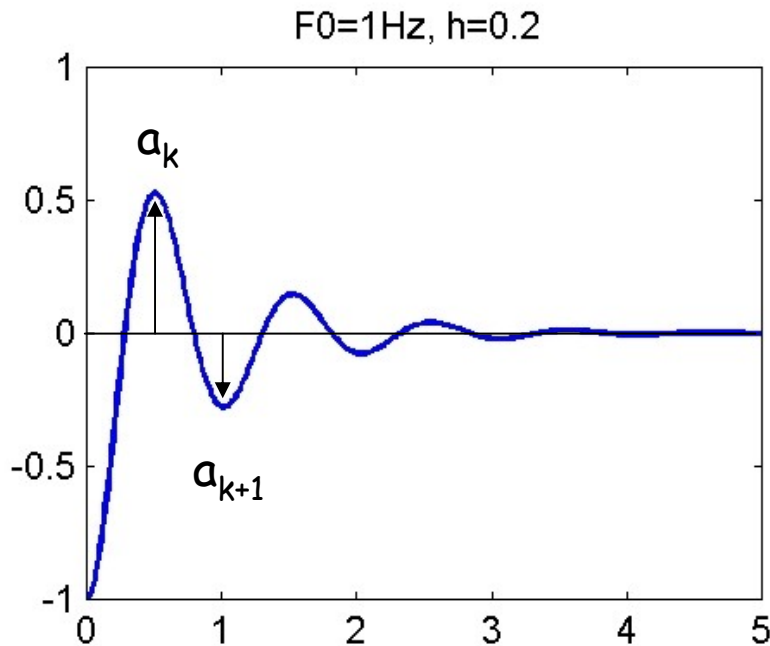
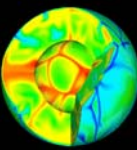


# Seismometer - Release Test





# Seismometer - Release Test

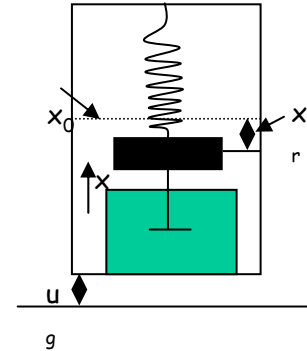


The damping coefficients can be determined from the amplitudes of consecutive extrema  $a_k$  and  $a_{k+1}$ . We need the logarithmic decrement  $\Lambda$

$$\Lambda = 2 \ln \left( \frac{a_k}{a_{k+1}} \right)$$

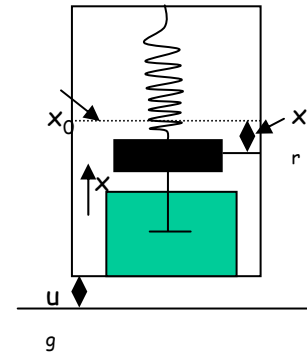
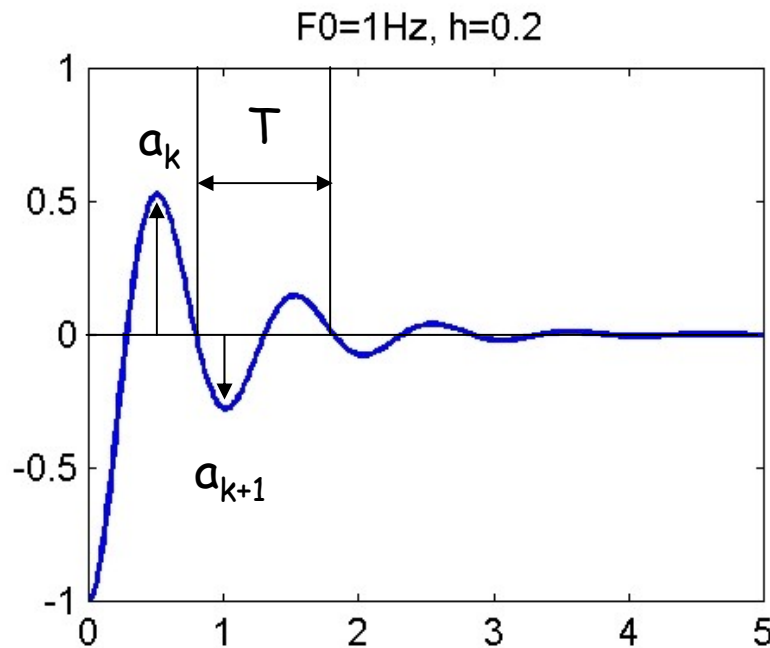
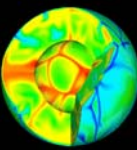
The damping constant  $h$  can then be determined through:

$$h = \frac{\Lambda}{\sqrt{4\pi^2 + \Lambda^2}}$$





# Seismometer - Frequency

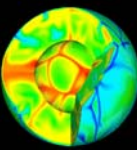


The period  $T$  with which the seismometer mass oscillates depends on  $h$  and (for  $h < 1$ ) is always larger than the period of the spring  $T_0$ :

$$T = \frac{T_0}{\sqrt{1-h^2}}$$

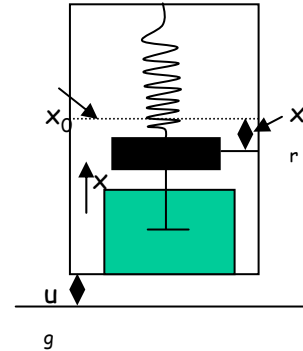


# Seismometer - Response Function



2. How does the seismometer amplify the ground motion? Is this amplification frequency dependent?

To answer this question we excite our seismometer with a monofrequent signal and record the response of the seismometer:



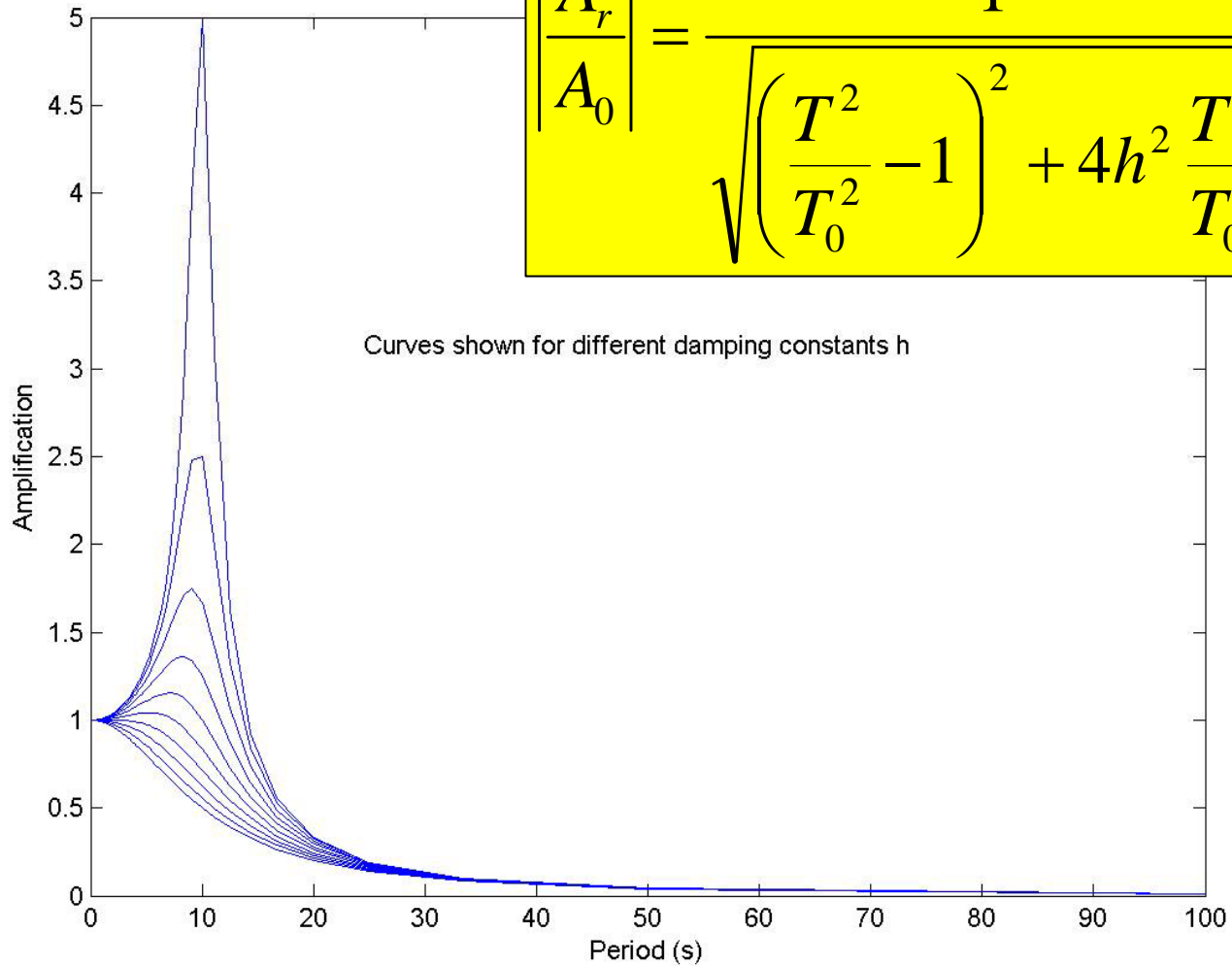
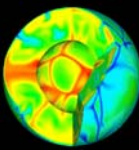
$$\ddot{x}_r(t) + h\omega_0\dot{x}_r(t) + \omega_0^2x_r(t) = \omega^2A_0e^{i\omega t}$$

the amplitude **response**  $A_r$  of the seismometer depends on the frequency of the seismometer  $\omega_0$ , the frequency of the excitation  $\omega$  and the damping constant  $h$ :

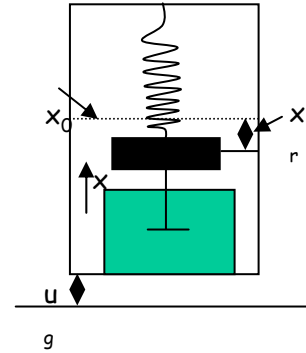
$$\left| \frac{A_r}{A_0} \right| = \frac{1}{\sqrt{\left( \frac{T^2}{T_0^2} - 1 \right)^2 + 4h^2 \frac{T^2}{T_0^2}}}$$



# Seismometer - Response Function



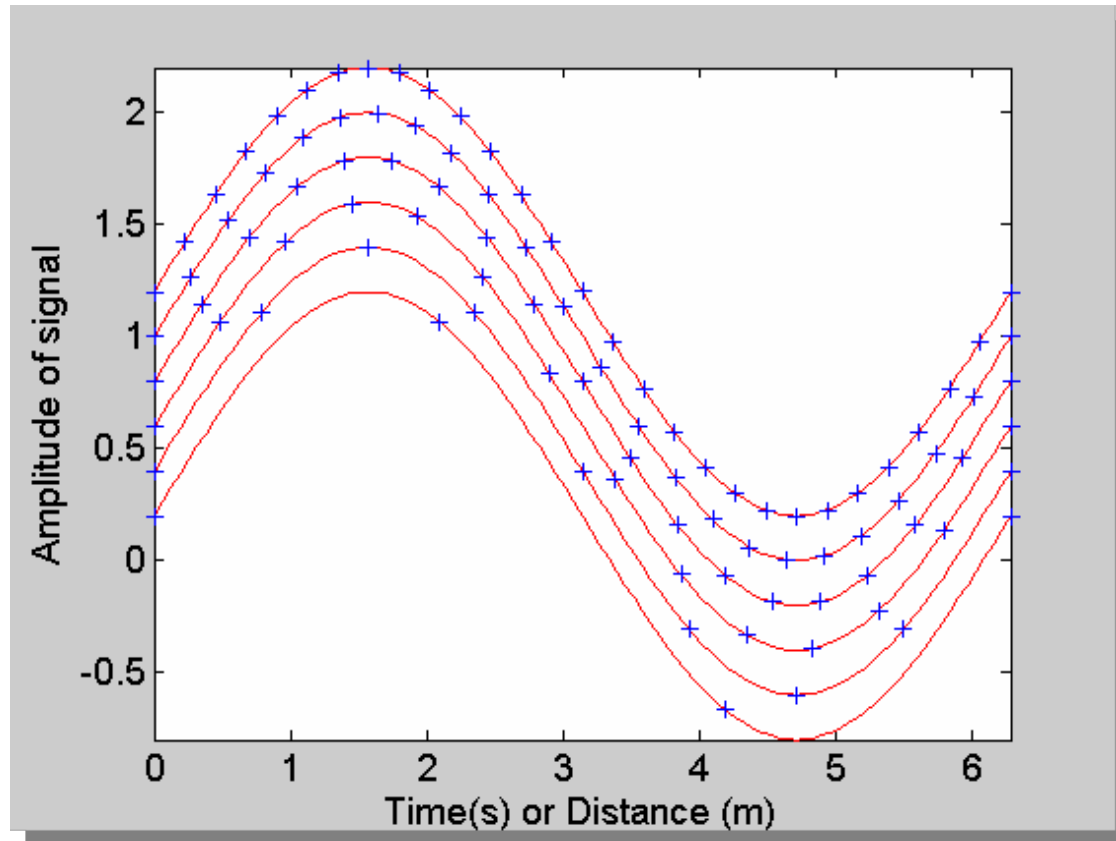
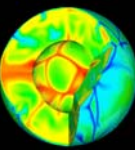
$$\left| \frac{A_r}{A_0} \right| = \frac{1}{\sqrt{\left( \frac{T^2}{T_0^2} - 1 \right)^2 + 4h^2 \frac{T^2}{T_0^2}}}$$







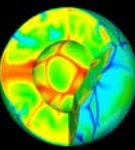
# Sampling rate



**Sampling frequency, sampling rate** is the number of sampling points per unit distance or unit time. Examples?



# Data volumes



Real numbers are usually described with 4 bytes (single precision) or 8 bytes (double precision). **One byte consists of 8 bits.** That means we can describe a number with 32 (64) bits. We need one switch (bit) for the sign (+/-)

-> 32 bits ->  $2^{31} = 2.147483648000000e+009$  (Matlab output)  
-> 64 bits ->  $2^{63} = 9.223372036854776e+018$  (Matlab output)  
(amount of different numbers we can describe)

How much data do we collect in a typical seismic experiment?

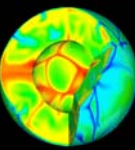
Relevant parameters:

- Sampling rate 1000 Hz, 3 components
- Seismogram length 5 seconds
- 200 Seismometers, receivers, 50 profiles
- 50 different source locations
- Single precision accuracy

How much (T/G/M/k-)bytes do we end up with? What about compression?



# (Relative) Dynamic range



What is the precision of the sampling of our physical signal in amplitude?

**Dynamic range**: the ratio between largest measurable amplitude  $A_{\max}$  to the smallest measurable amplitude  $A_{\min}$ .

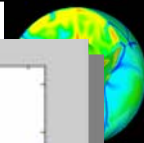
The unit is Decibel (dB) and is defined as the ratio of two power values (and power is proportional to amplitude square)

In terms of amplitudes

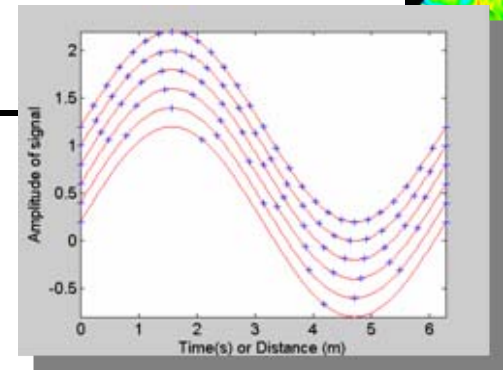
$$\text{Dynamic range} = 20 \log_{10}(A_{\max}/A_{\min}) \text{ dB}$$

Example: with 1024 units of amplitude ( $A_{\min}=1$ ,  $A_{\max}=1024$ )

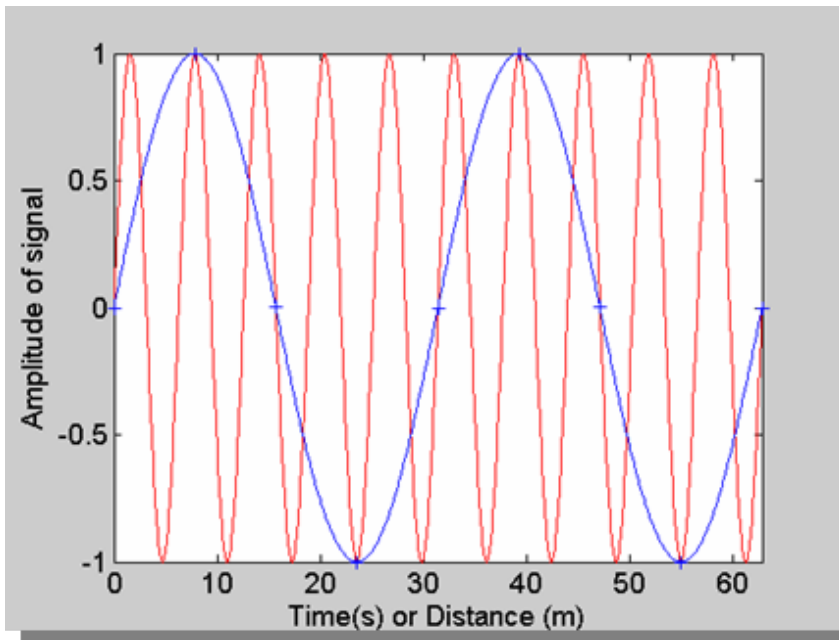
$$20 \log_{10}(1024/1) \text{ dB} \approx 60 \text{ dB}$$



# Nyquist Frequency (Wavenumber, Interval)



The frequency **half of the sampling rate**  $dt$  is called the **Nyquist frequency**  $f_N=1/(2dt)$ . The distortion of a physical signal higher than the Nyquist frequency is called **aliasing**.

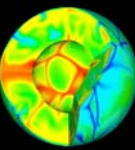


The frequency of the **physical signal** is  $> f_N$  is sampled with (+) leading to the erroneous **blue** oscillation.

What happens in space?  
How can we avoid aliasing?



# A cattle grid

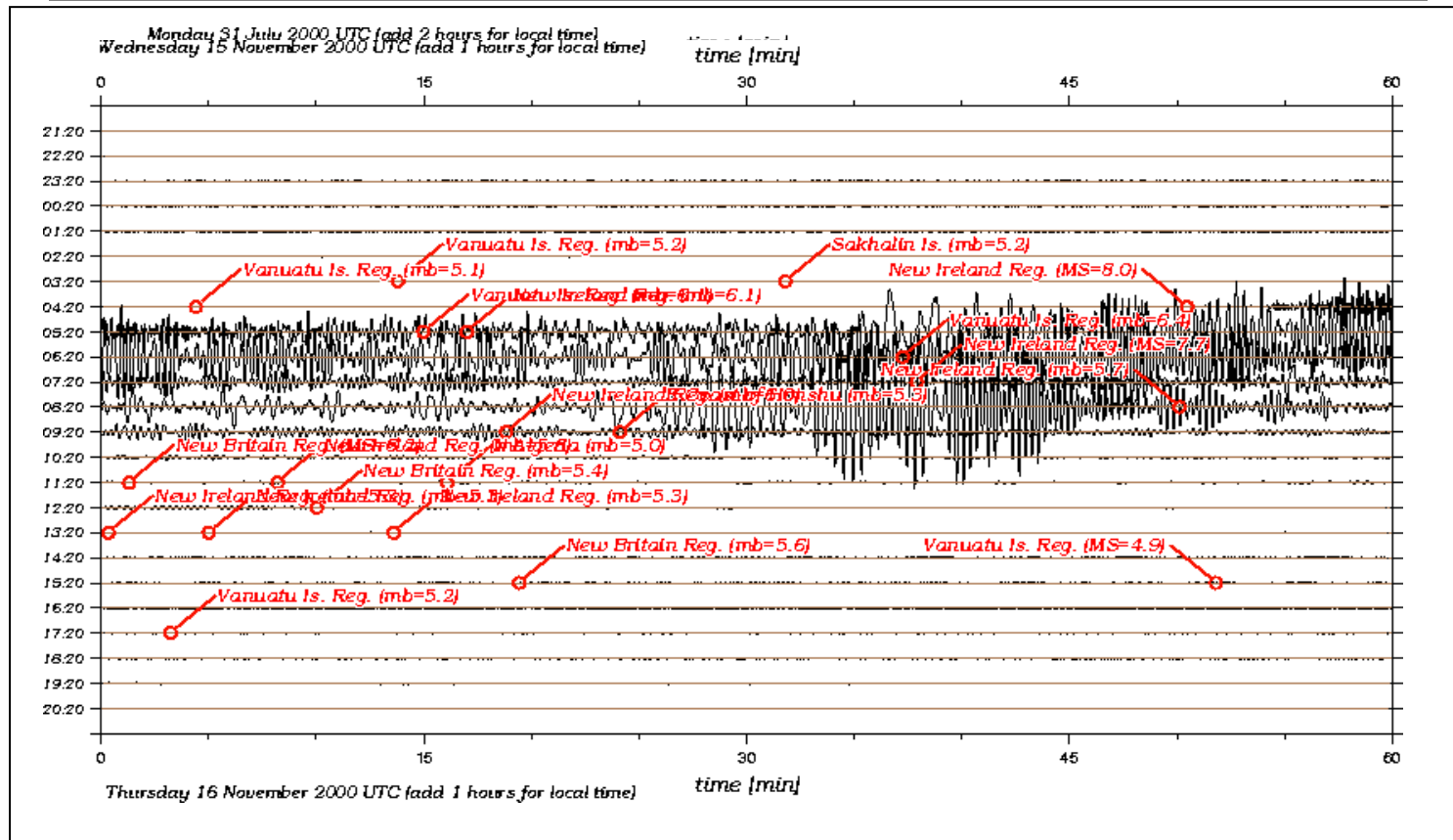




# Signal and Noise

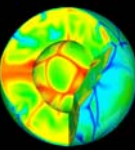


Almost all signals contain **noise**. The **signal-to-noise ratio** is an important concept to consider in all geophysical experiments. Can you give examples of noise in the various methods?





# Discrete Convolution



**Convolution** is the mathematical description of the change of waveform shape after passage through a filter (system).

There is a special mathematical symbol for convolution (\*):

$$y(t) = g(t) * f(t)$$

Here the impulse response function  $g$  is convolved with the input signal  $f$ .  $g$  is also named the „Green's function“

$$y_k = \sum_{i=0}^m g_i f_{k-i}$$

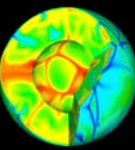
$$k = 0, 1, 2, \dots, m + n$$

$$g_i \quad i = 0, 1, 2, \dots, m$$

$$f_j \quad j = 0, 1, 2, \dots, n$$



# Convolution Example (Matlab)



```
>> x
x =
    0    0    1    0

>> y
y =
    1    2    1

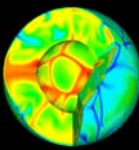
>> conv(x,y)
ans =
    0    0    1    2    1    0
```

Impulse response

System input

System output



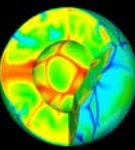


# Convolution Example (pictorial)

x	„Faltung“	y	x*y
	0 1 0 0	1 2 1	0
	0 1 0 0	1 2 1	0
	0 1 0 0	1 2 1	1
	0 1 0 0	1 2 1	2
	0 1 0 0	1 2 1	1
1 2 1	0 1 0 0	1 2 1	0

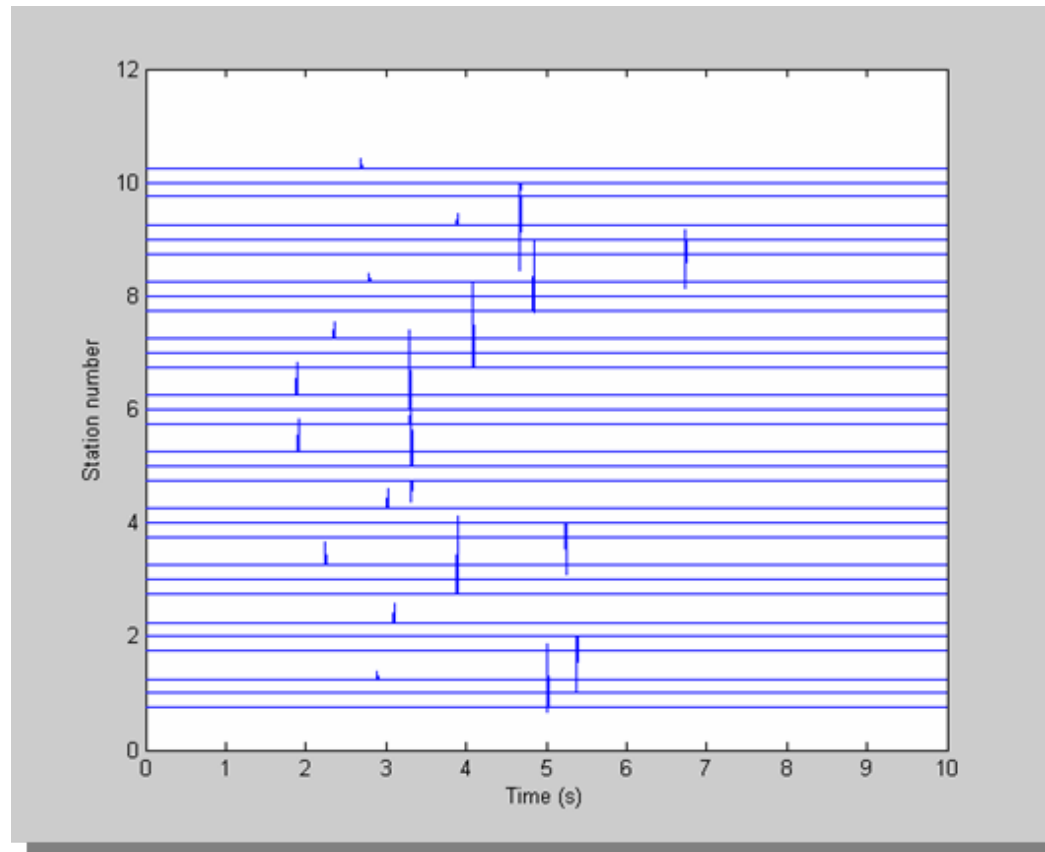


# Deconvolution



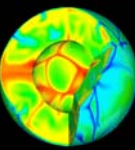
**Deconvolution** is the inverse operation to **convolution**.

When is **deconvolution** useful?





# Digital Filtering



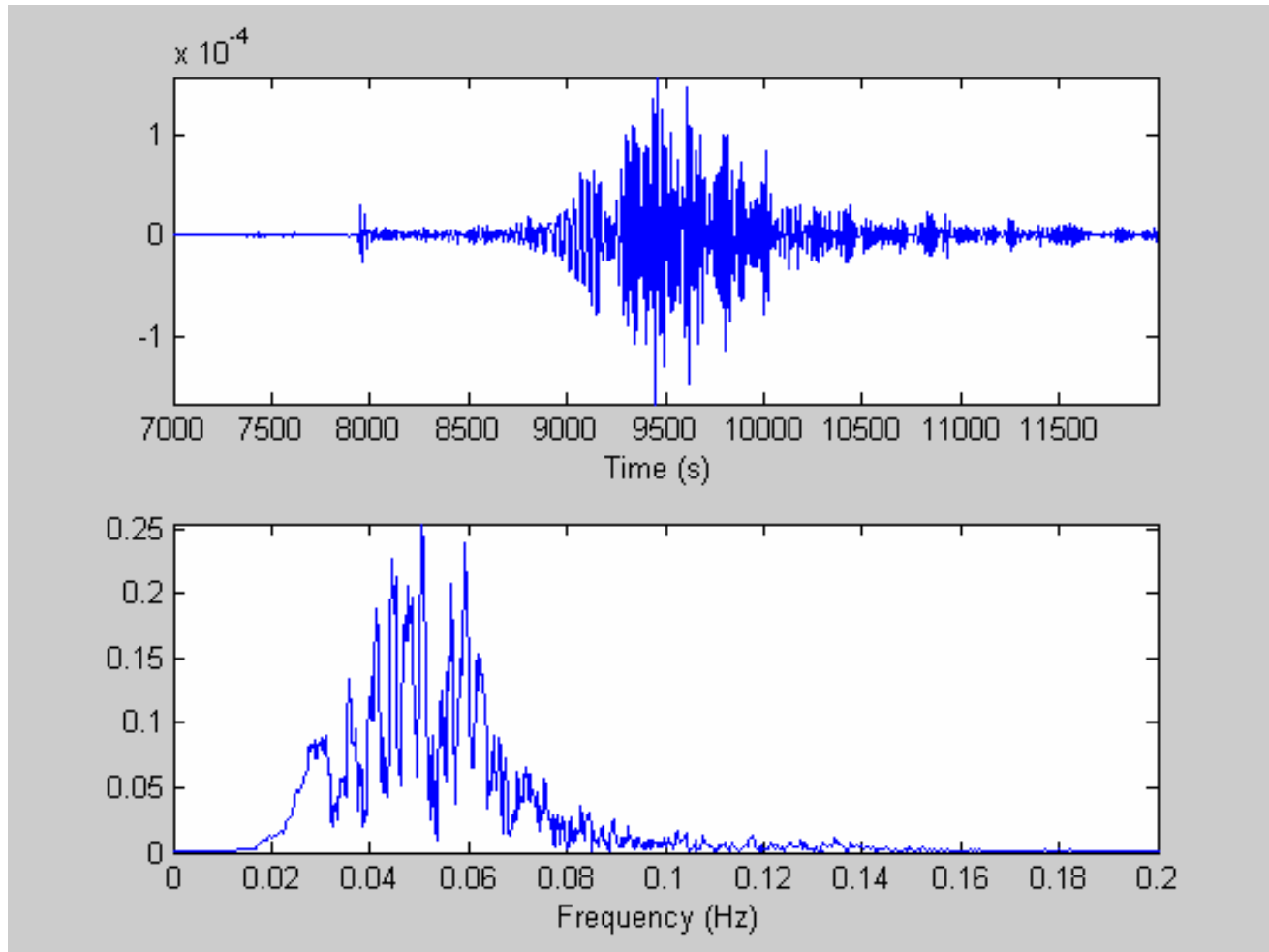
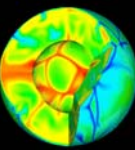
Often a recorded signal contains a lot of information that we are not interested in (noise). To get rid of this noise we can apply a **filter in the frequency domain**.

The most important filters are:

- **High pass:** cuts out low frequencies
- **Low pass:** cuts out high frequencies
- **Band pass:** cuts out both high and low frequencies and leaves a band of frequencies
- **Band reject:** cuts out certain frequency band and leaves all other frequencies

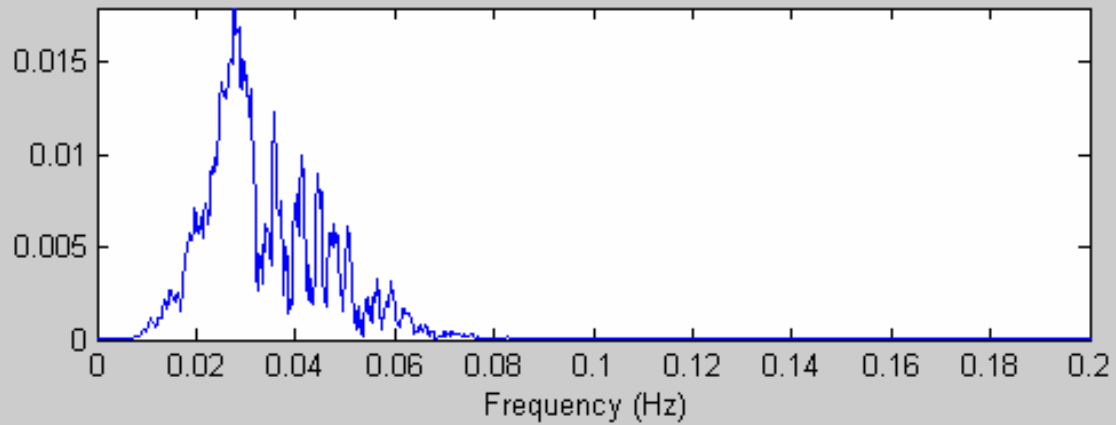
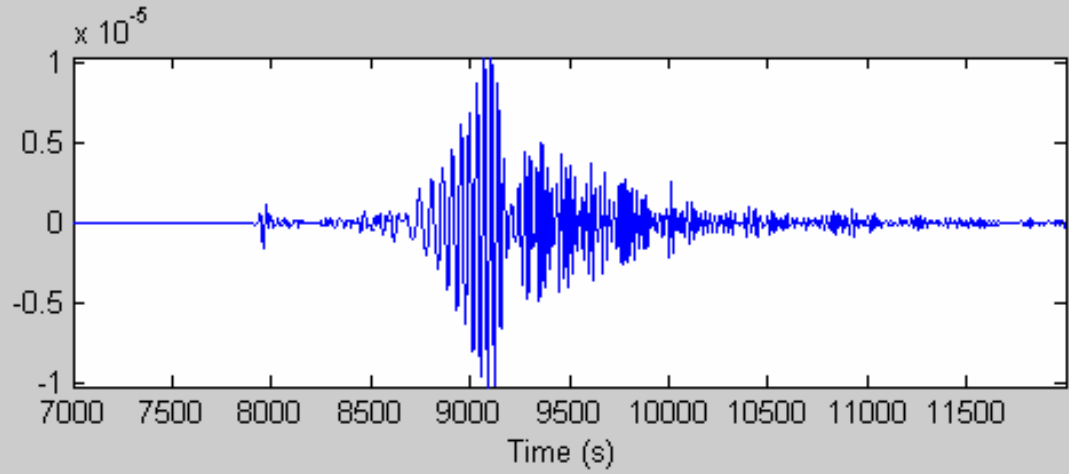
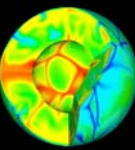


# Digital Filtering



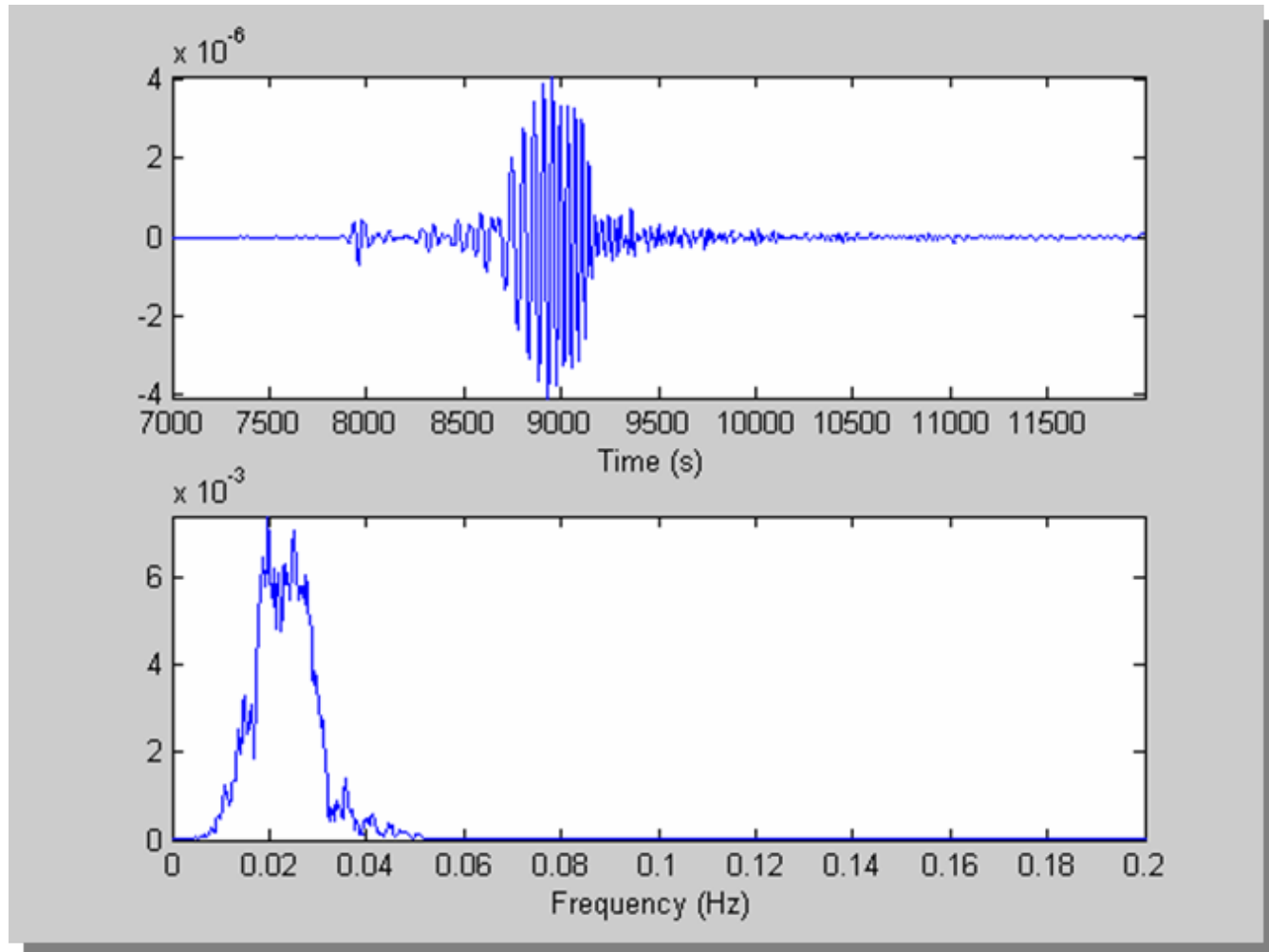
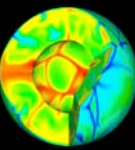


# Low-pass filtering



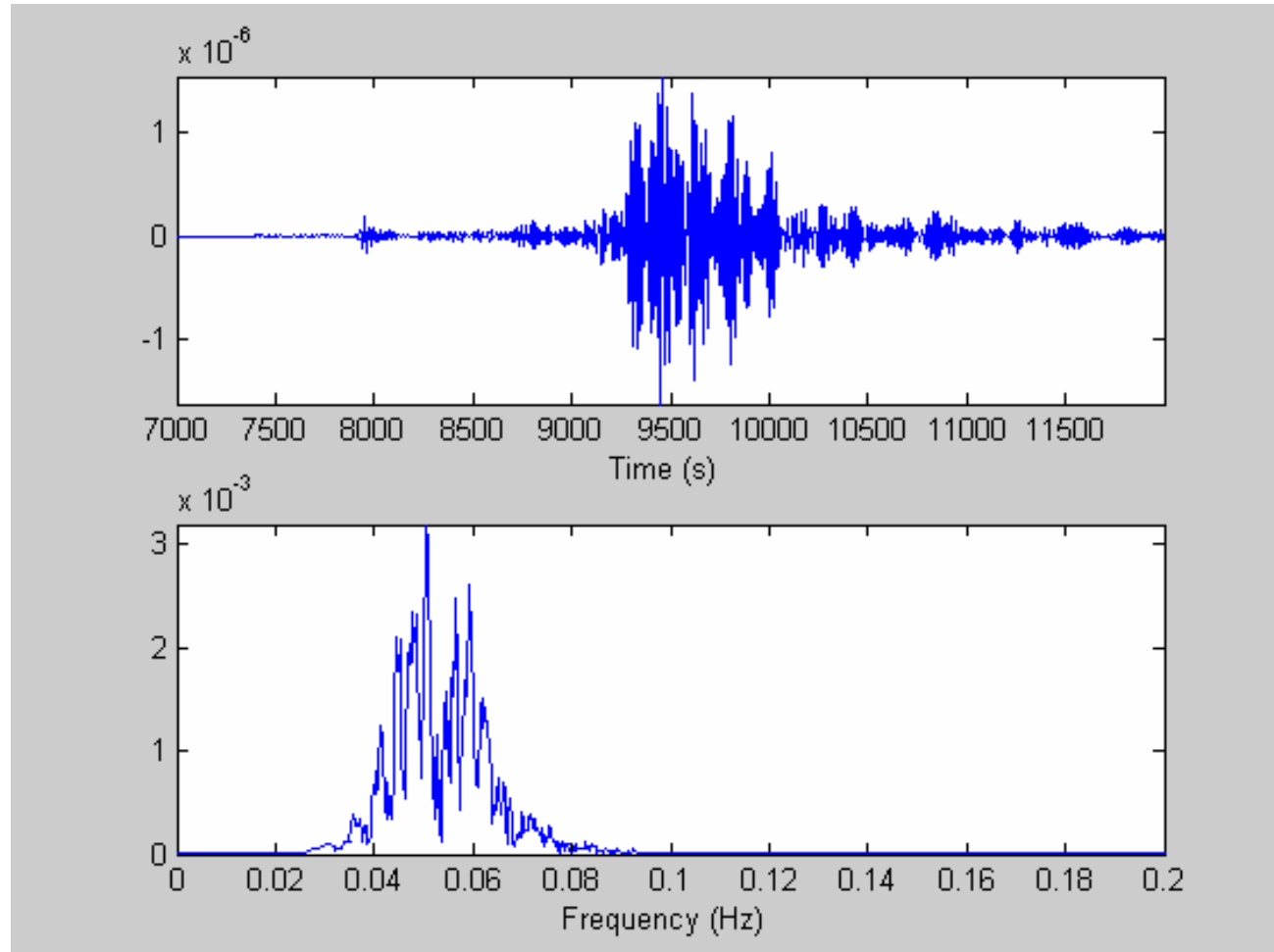
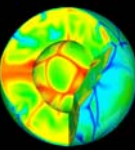


# Lowpass filtering



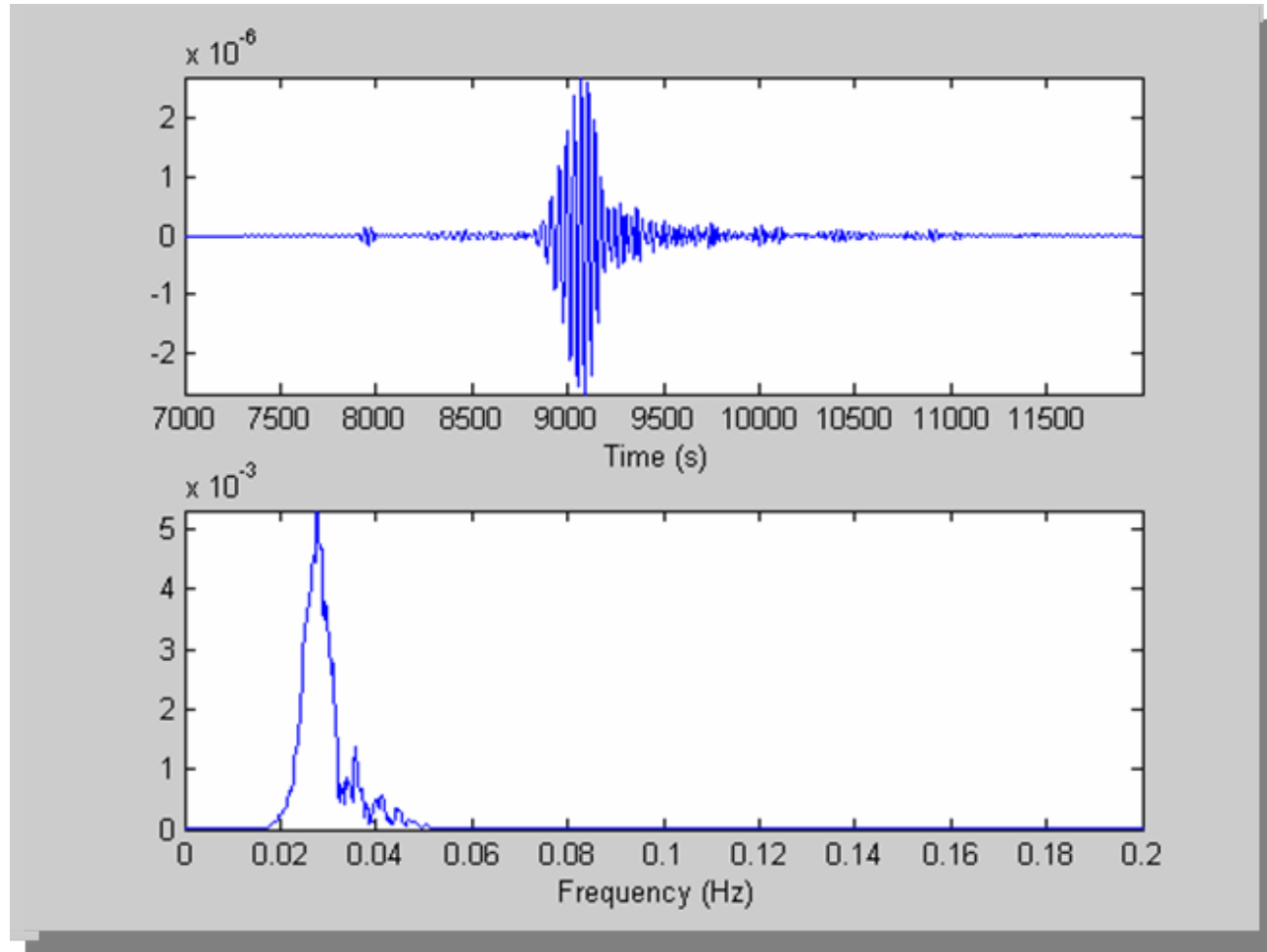
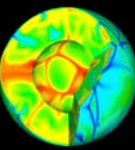


# High-pass filter





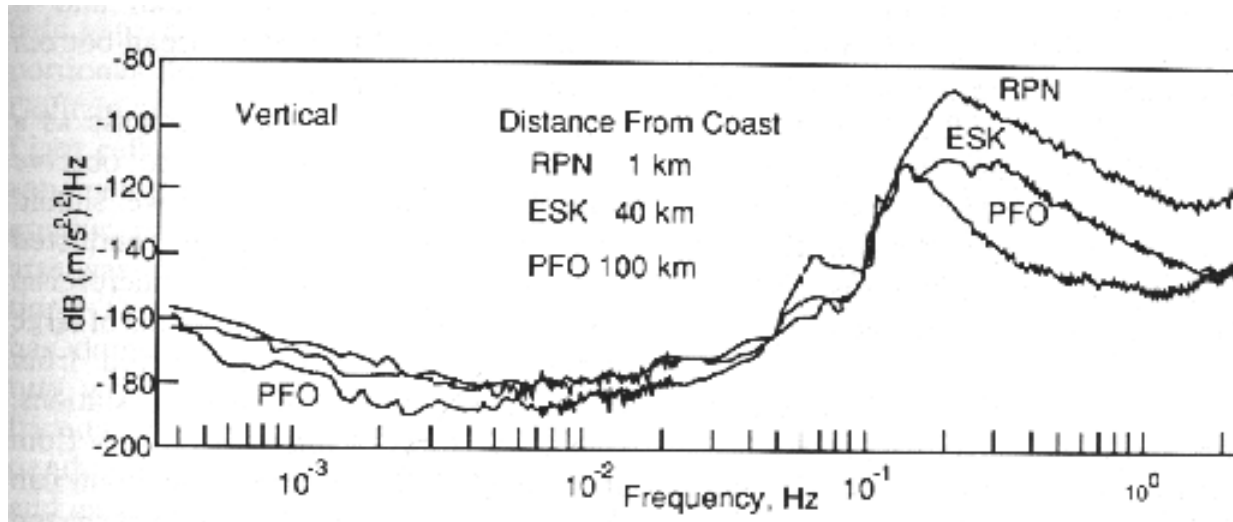
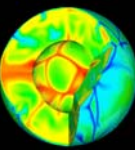
# Band-pass filter







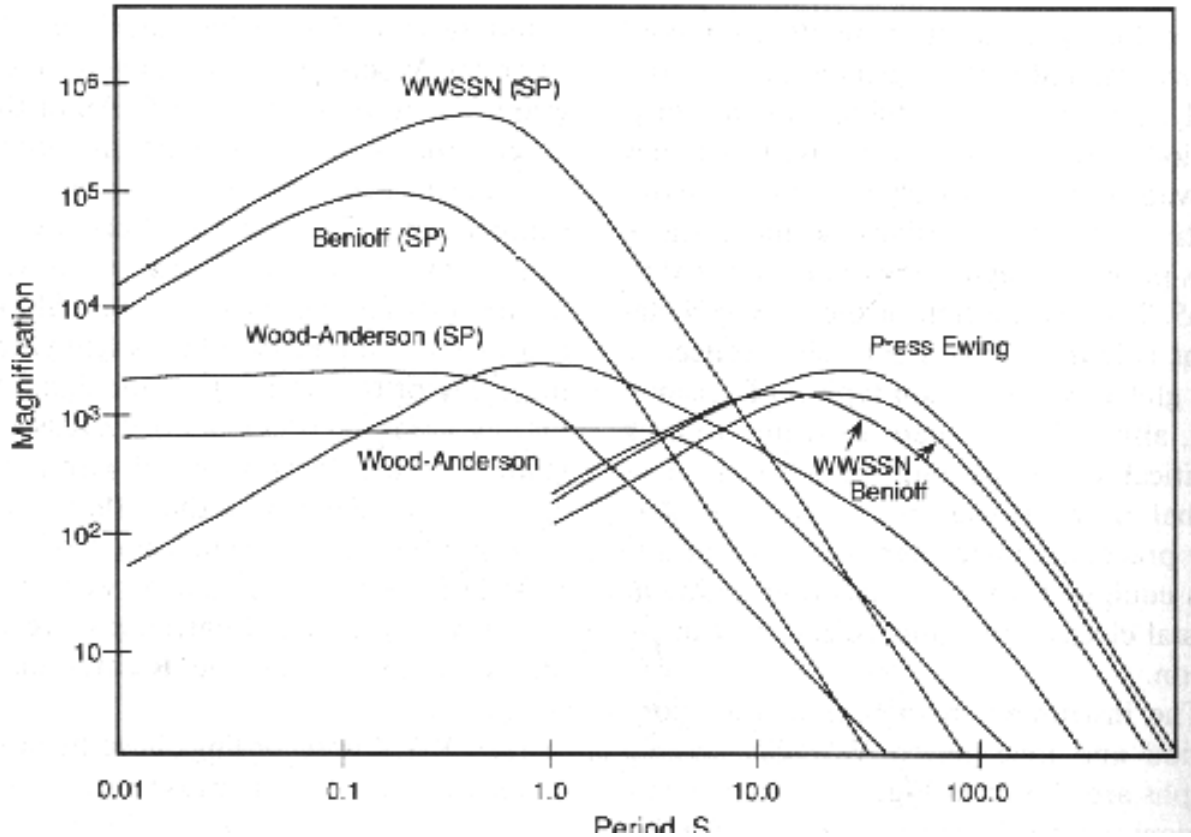
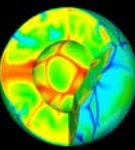
# Seismic Noise



Observed seismic noise as a function of frequency (power spectrum). Note the peak at 0.2 Hz and decrease as a distant from coast.



# Instrument Filters





# Time Scales in Seismology

